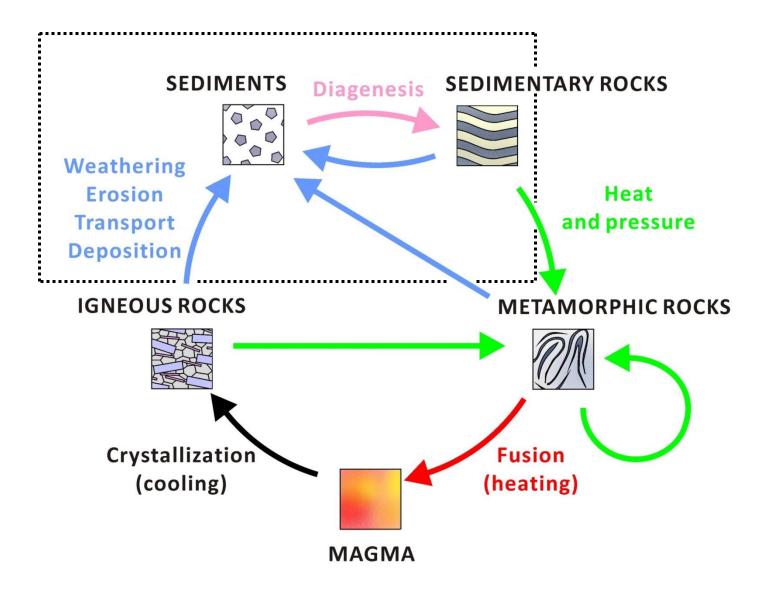
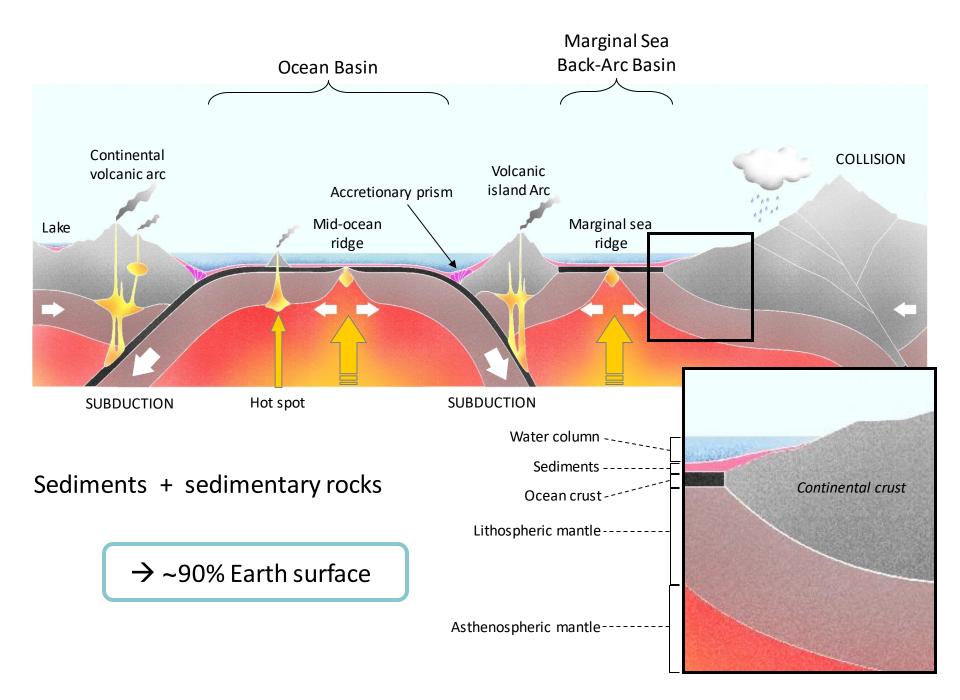
Sedimentary rocks





★ What are sediments?

Sediments are	Examples of sedimentary rocks
1. Solid particles	↓ ↓
Rocks/minerals Ex.: Quartz sand	Sandstone
 Hard parts of organisms (biominerals) Ex.: Foraminiferal sand (CaCO₃) 	Foraminiferal limestone
Organic material Ex.: Plant debris (peat)	
2. Dissolved chemicals	
 Abiotic precipitation Ex.: Na⁺, K⁺, Cl⁻ 	Evaporite: NaCl, KCl
 Biotic precipitation Ex.: Corals (CaCO₃) 	Reeflimestone



www.pitt.edu



Sandstone

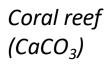


www.microimaging.ca

Foraminiferal sand

Quartz sand

Peat





Reef limestone

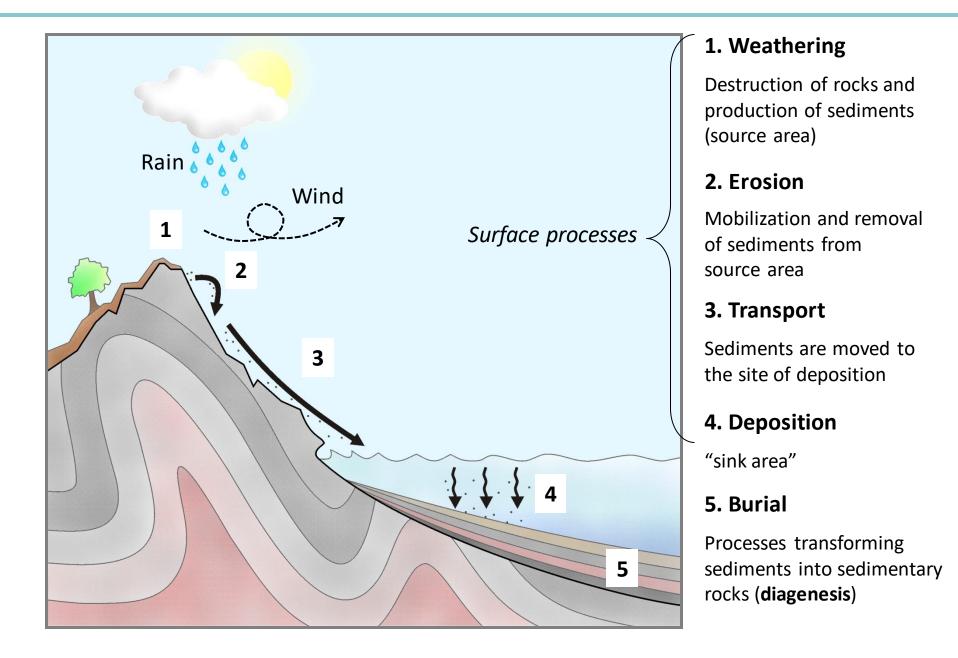
Coal



Foraminiferal limestone

Wikipedia

***** Formation of sediments and sedimentary rocks



O Weathering processes

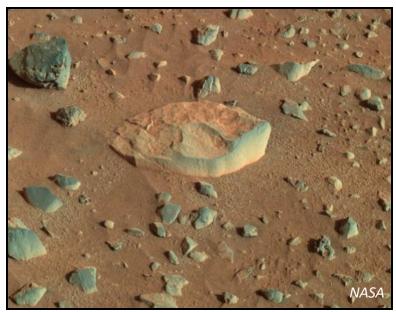
Physical weathering

- Mechanical weathering by wind (1), water (2) and ice (3, 4)
- Biophysical weathering e.g. root wedging (5)
- (1) Weathering by wind

Moroccan desert pavement

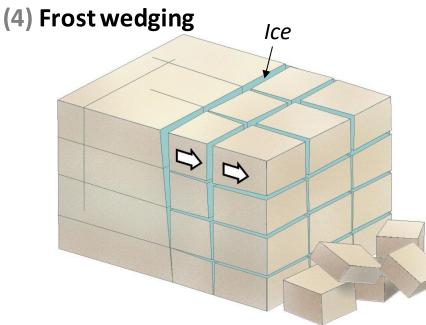


Mars



(2) Weathering by waves





(3) Weathering by glaciers



(5) Root wedging





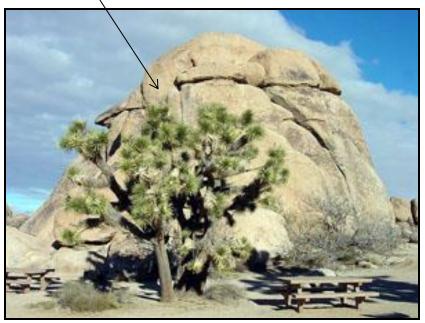
Chemical weathering

- Dissolution of minerals (mainly CaCO₃) by mildly acidic water (1)
- Biotic mineral dissolution e.g. microbes, lichen, clionid sponge (2,3)
- (1) Weathering/dissolution of carbonates (karst, e.g. caves)



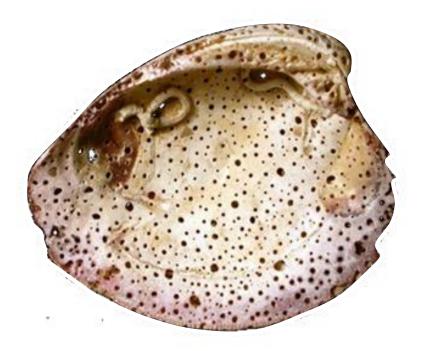
Carsten Peter, National Geographic

Weathering/dissolution of silicates Body of granite rounded by weathering and erosion \

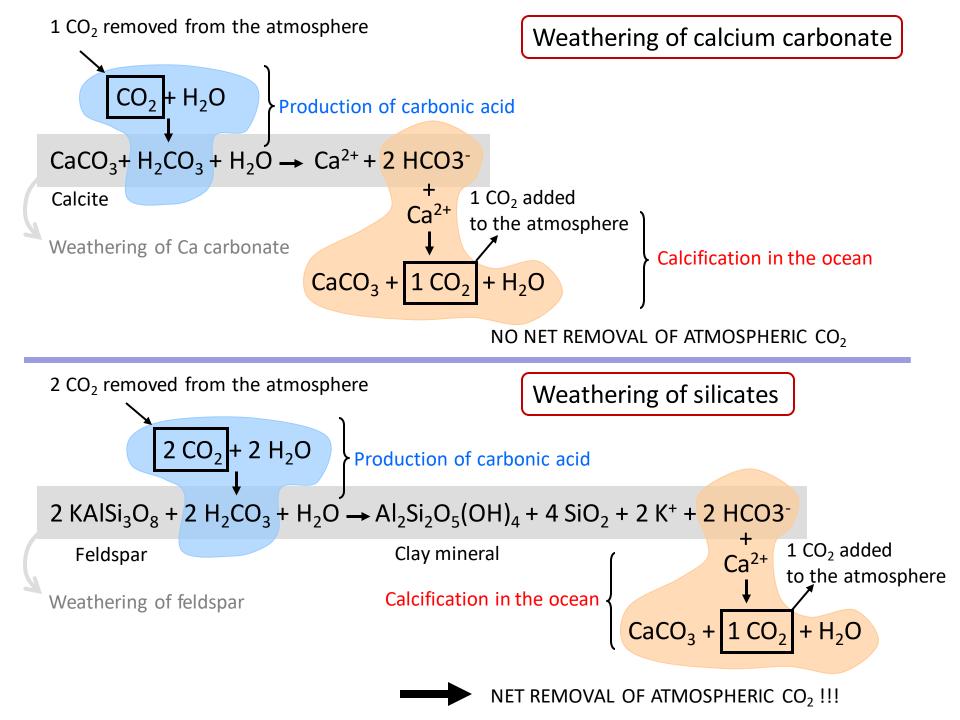


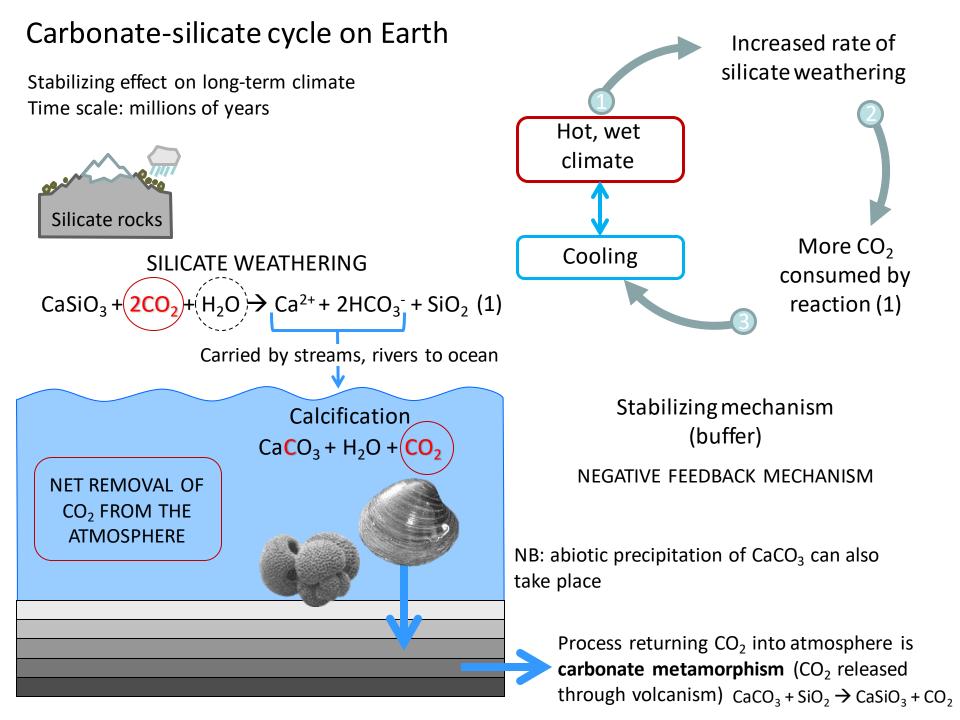
(2) Boreholes of clionid sponge

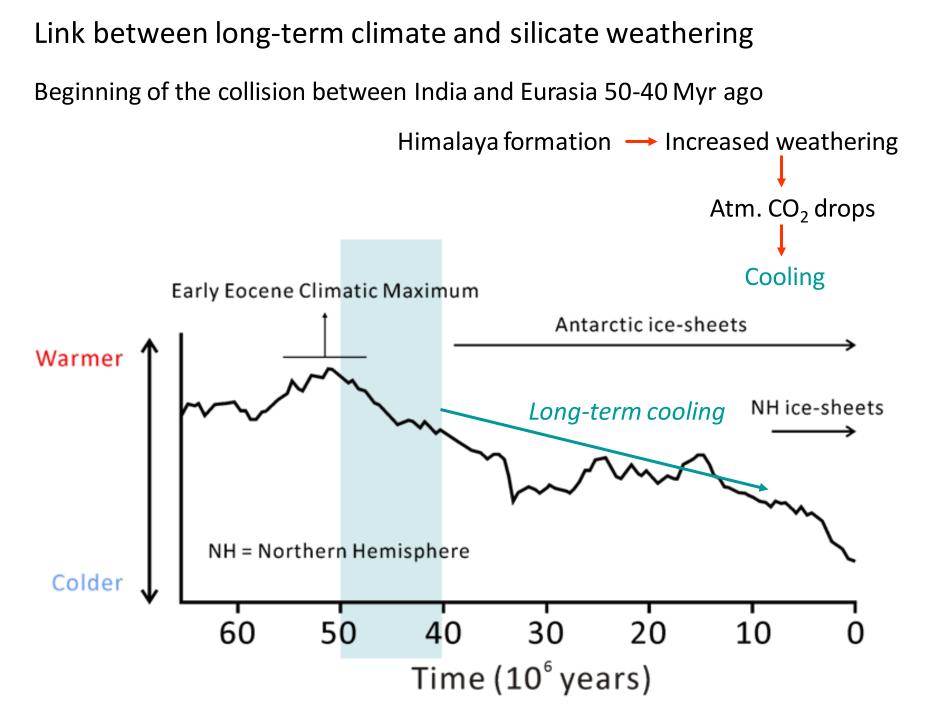
(3) Boreholes of bivalve



Mark A. Wilson (Dep. of Geology, College of Wooster) Biolib







4KAlSi₃O₈ + 4H⁺ + 2H₂O \rightarrow 4K⁺ + Al₄Si₄O₁₀(OH)₈ + 8SiO₂ Orthoclase Kaolinite Quartz

Hydrolysis of granite

Remobilized and transported by rain water and deposited in depressions



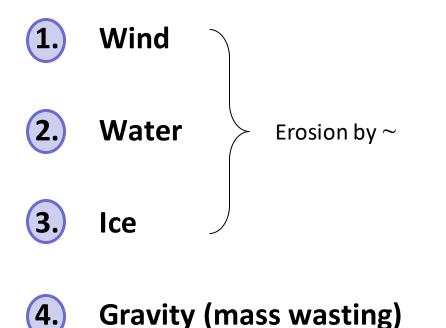
Kaolin quarry (Japan) http://www.eacrh.net/ojs/index.p hp/crossroads/article/view/14/Vo I3_Seyock_html

NB: Kaolinite is primarily used in the paper industry (paper coating)

O Erosion and transport

"As soon as a rock particle (loosened by one of the two weathering processes) moves, we call it **erosion** or **mass wasting**. Mass wasting is simply movement down slope due to gravity. Rock falls, slumps, and debris flows are all examples of mass wasting. We call it erosion if the rock particle is moved by some flowing agent such as air, water or ice."

From USGS





Great Sand Dunes National Park (Colorado, USA)



Walter Meayers Edwards, National Geographic



Idaho (USA)

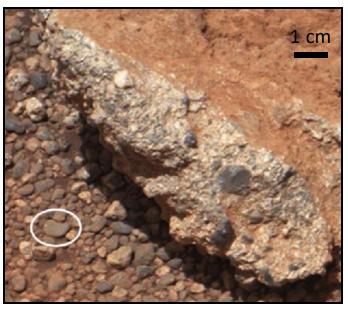


Michael Melford, National Geographic



Whirlwind (dust devil) on Mars

NASAAncient fluvial deposit on Mars NASA



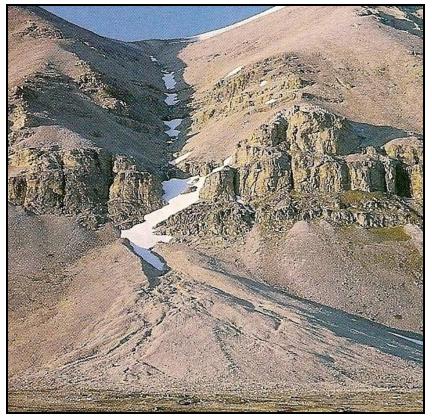




Glacier in British Columbia (Canada)



Debris cone (Spitzberg, Norway)

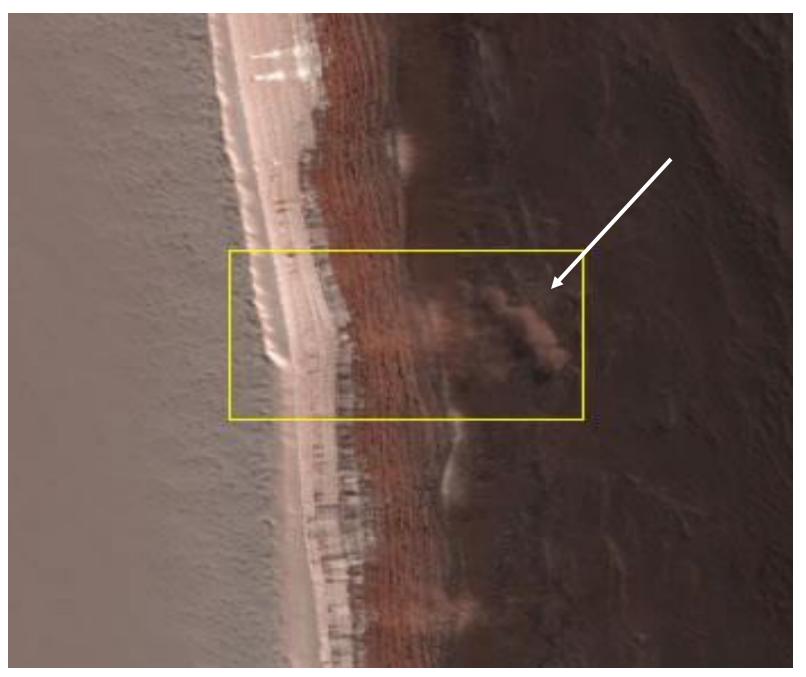


Chenuet (1993)

Sarah Leen, National Geographic

Glacial grooves formed during the last glaciation (Kelleys Island, Ohio)

Wikipedia



O Sediment deposition

1. Water/wind

As **wind/water current** decreases, it can no longer keep the largest particles suspended.

The stronger the current, the larger the particles it can carry:

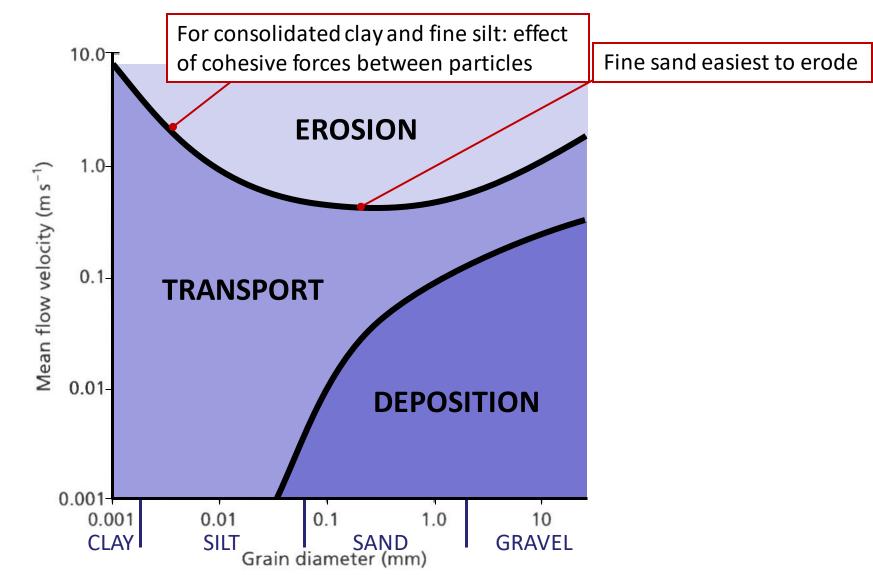
<u>Strong currents</u> (>50 cm/s): carry gravels (>2 mm) and smaller particles <u>Moderately strong currents</u> (20-50 cm/s): carry sand grains (62.5 μm-2 mm) and smaller particles <u>Weak currents</u> (<20 cm/s): carry silt and clay particles (mud; <62.5 μm)

2. Ice

Sediments are deposited as ice melts and retreats.

3. Gravity

Deposition is controlled by topography (slope steepness) and the nature of sediments (size, shape)



HJÜLSTROM DIAGRAM Deduced experimentally (for sediments transported by water) **Fig. 2.14** Schematic representation of relationship between current velocity and sediment erosion, transport and deposition (Hjülstrom's diagram, deduced experimentally from flows of 1 m depth). Note that sediment may continue to be transported after the current velocity has fallen below the level at which it was initially eroded. (simplified)

Estuary mud flats

http://gravelbeach.blogspot.com/2016/10/mulranny-beach.html

https://geologicalintroduction.baffl.co.uk/?attachment_id=453

Glacial erratic



Glacial till (moraine) – coarse unsorted sediment in fine-grained (clay) matrix



Glacial striation



Glacial valley



O Burial and diagenesis: sedimentary rock formation

- Burial: process by which sediments are buried under new layers of sediments → increase in temperature and pressure
- **Diagenesis**: set of physical and chemical changes affecting **sediments** after they are buried.

Diagenetic processes leading to **lithification**:

- Compaction (due to burial)
- Cementation

Decrease in porosity (% of rock's volume consisting of open space / pores)

Transformation of soft sediments into hard sedimentary rocks

= Lithification

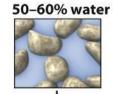
- 1 Sediments are buried, compacted, and lithified at shallow depths in Earth's crust.
- 2 Diagenesis includes the processes —physical and chemical—that change sediments to sedimentary rocks.

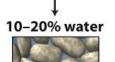
Compaction

Cementation

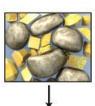
Lithification

Compaction by burial squeezes out water. Precipitation or addition of new minerals cements sediment particles.

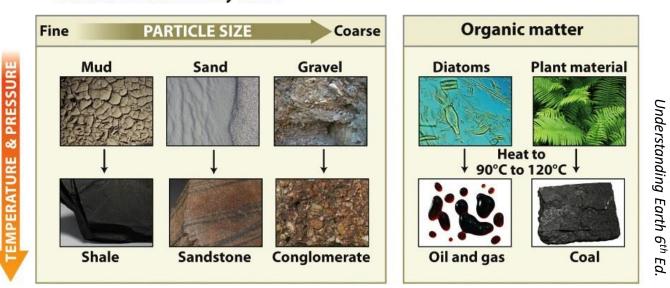








3 Different sediments result in different sedimentary rocks.



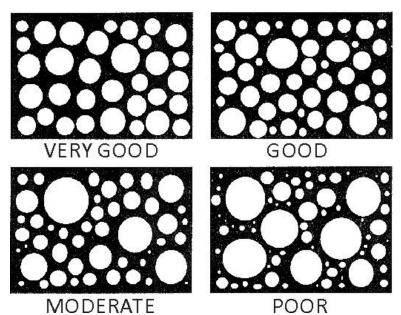
★ Properties of sediments and sedimentary rocks

• Grain size → Influenced by wind/water velocity



(Shale breaks along stratification planes, mudstone does not)

• Sorting

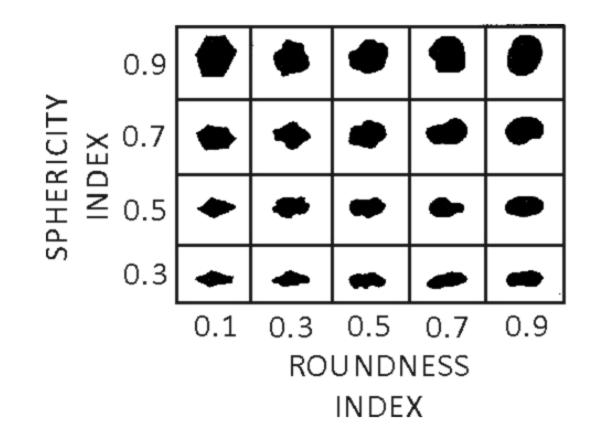


Good sorting indicates a transport agent of constant strength

Poor sorting indicates a transport agent of variable strength

• Grain morphology

The degree of abrasion (roundness) depends on the distance of transport.

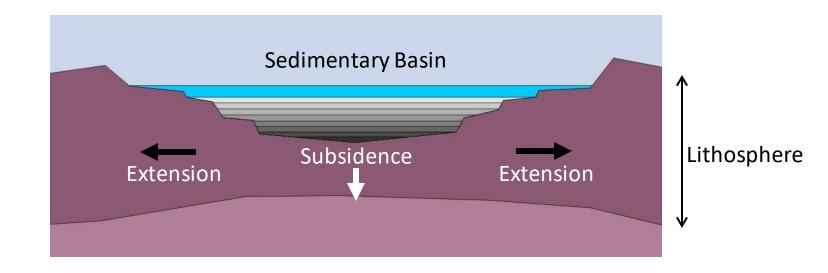


***** Sedimentary basins and sedimentary environments

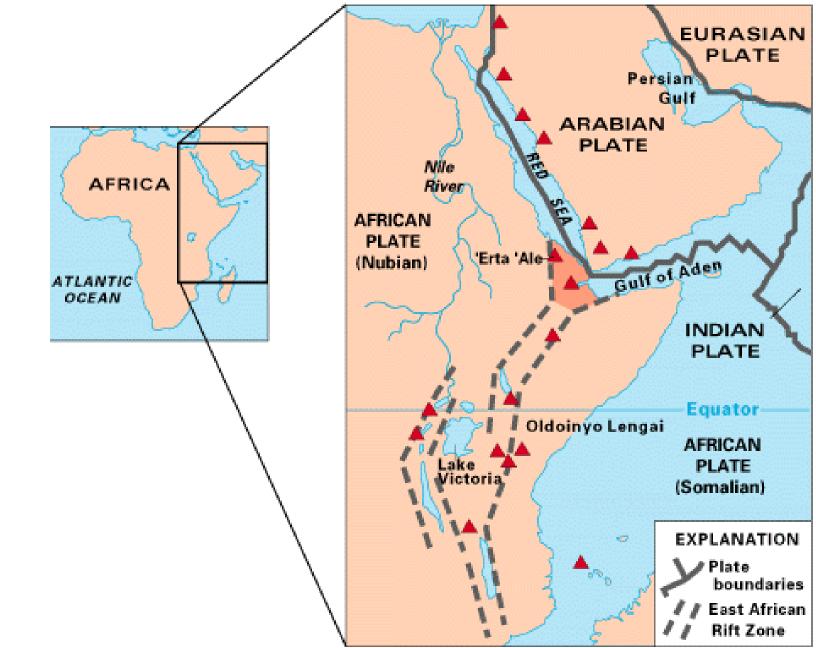
- Sediments tend to accumulate in **depressions**.
- Large depressions are formed by **subsidence**.

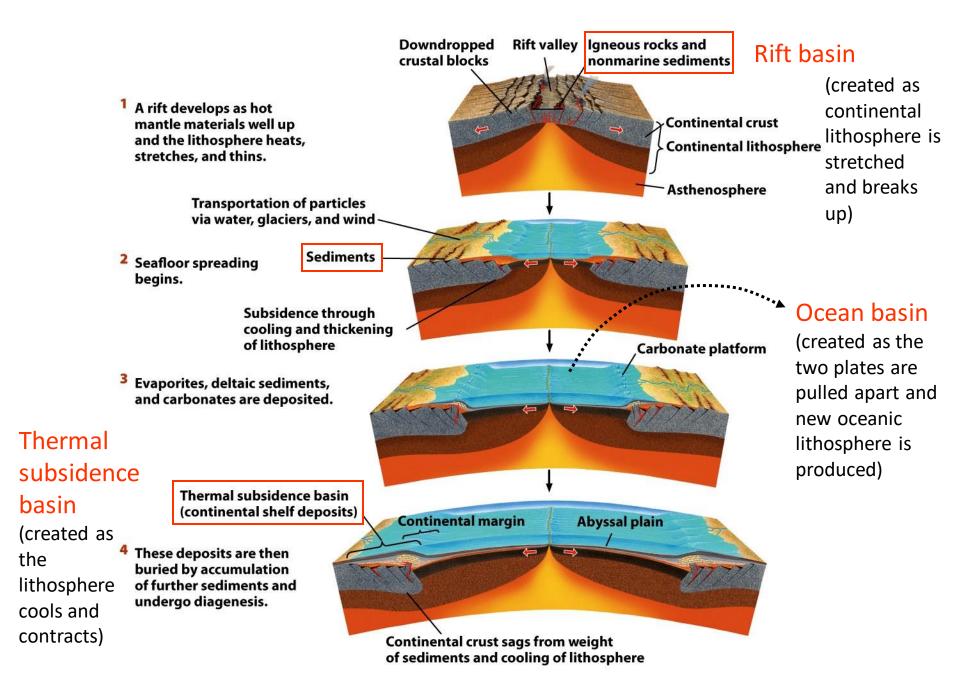
Subsidence is the process by which a broad area of the crust sinks (subsides) relative to the surrounding crust. It is mainly due to **tectonic deformation** of the lithosphere (stretching) and accentuated by the **weight of sediments**.

• Regions characterized by thick accumulations of sediments and sedimentary rocks are called **sedimentary basins**.

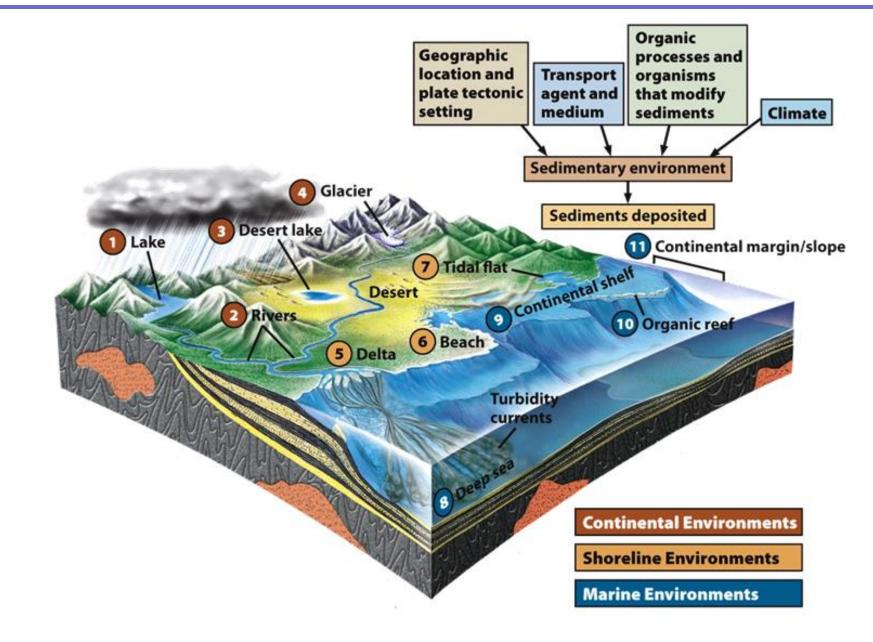


East African Rift





Sedimentary environments



Continental environments





Shoreline environments

Deltas

Tidal flats

Marine environments

Coral reefs

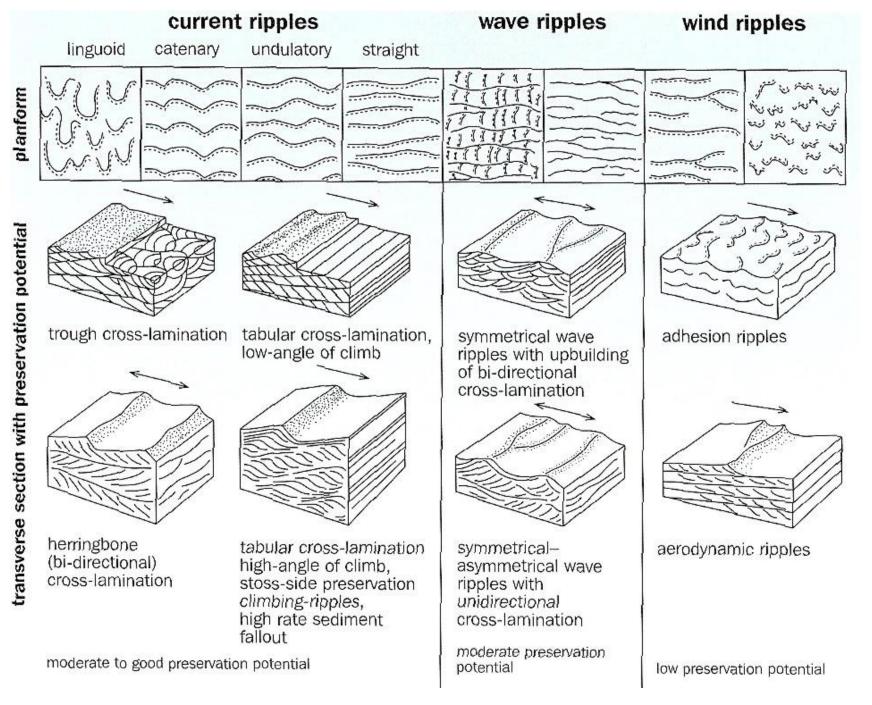
dip.

Deep sea

***** Examples of sedimentary structures

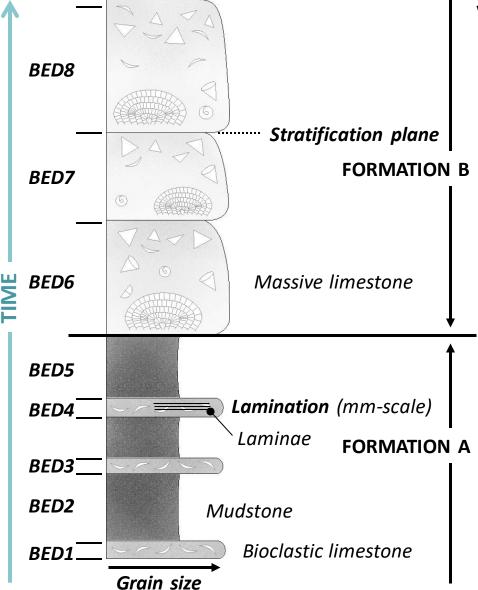
1. Ripple marks



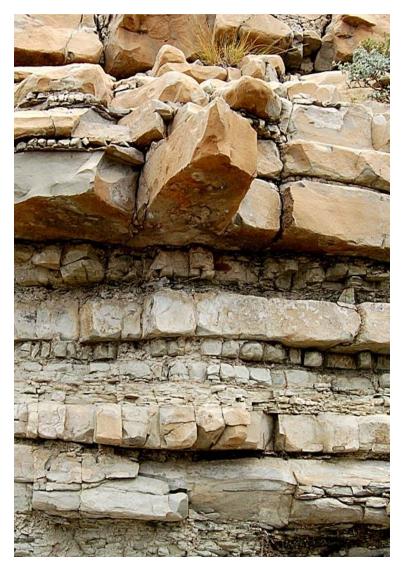


From Stow (2005)





Stratification plane: separation between two beds (originally horizontal if sediments were deposited as flat-lying layers or inclined if they were deposited on a slope)



www.edupic.net

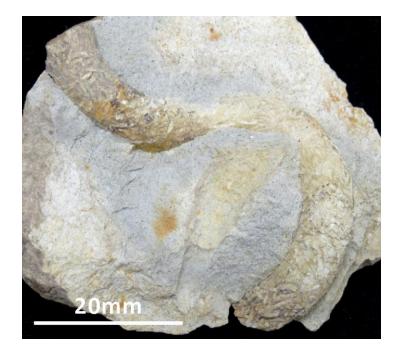


Bioturbation (disturbance of soils and sediments by animals or plants)

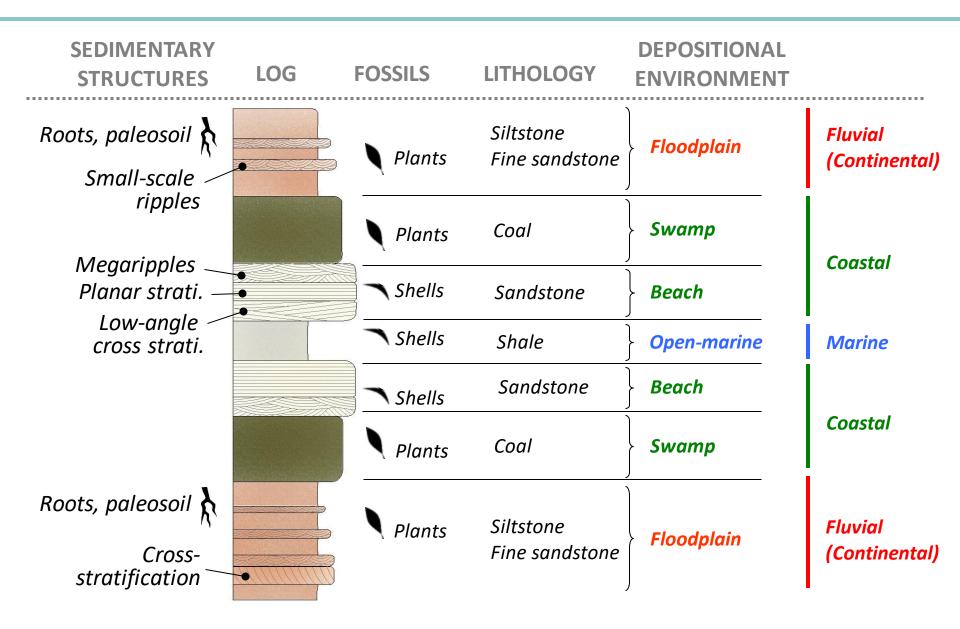
Large burrows in volcanic tuff (Holocene, Japan)



Small burrow (Paleozoic, Belgium)

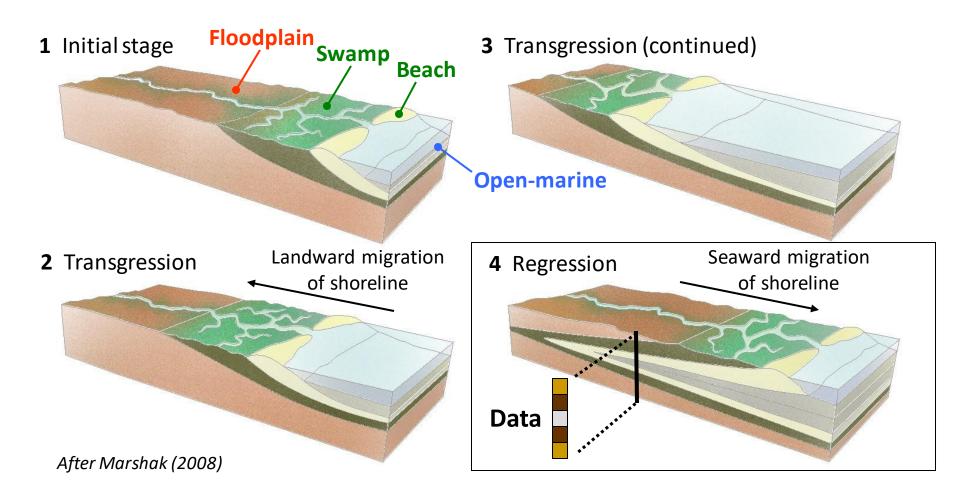


★ Reconstruction of past sea level and environmental changes



After Marshak (2008)

Data interpretation



Transgression = sea level rises and shoreline moves landward **Regression** = sea level falls and shoreline moves seaward

★ Types of sedimentary rocks

1. Siliciclastic sedimentary rocks

 Derived from accumulation of mineral fragments and/or lithic (rock) fragments composed mainly of silicate minerals

2. Biochemical sedimentary rocks

 Derived from precipitation (direct or indirect) of a mineral by organisms (most commonly CaCO₃ or SiO₂)

3. Chemical sedimentary rocks

 Derived from abiotic precipitation of minerals in a saline pond, lake or embayment undergoing intense evaporation (→ EVAPORITES)

4. Organic sedimentary rocks

• Derived from accumulation and preservation of organic matter



• CLASSIFICATION BASED ON GRAIN SIZE:

COARSE-GRAINED Conglomerate







FINE-GRAINED Shale



CLASSIFICATION BASED ON GRAIN COMPOSITION

Example: sandstones

Lithic sandstone Rich in rock fragments Arkose Feldspar-rich Quartz arenite Pure quartz Graywacke

Space between sand grains filled with mud

Biochemical sedimentary rocks



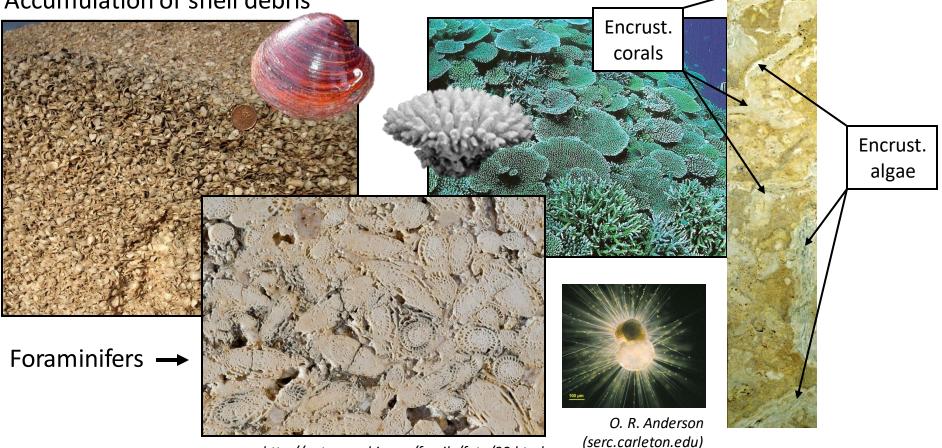
2.)

Direct precipitation of CaCO₃ by corals, mollusks, foraminifera, diatoms = LIMESTONES

SiO₂ by diatoms, radiolarians

Reef limestones

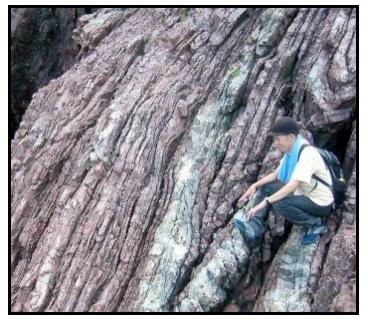
Bioclastic limestones Accumulation of shell debris

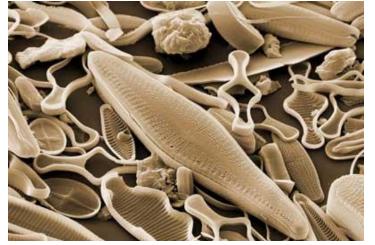


http://petrographica.ru/fossils/foto/90.html

CHERT = siliceous sedimentary rock (composed of silica)

Radiolarite (Japan)





www.radiolaria.org

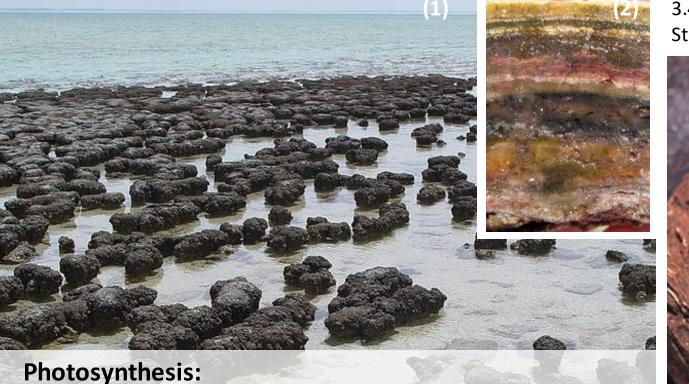
Like foraminifers and radiolarians, diatoms are single-celled organisms. However, diatoms can sometimes form colonies of attached individuals.

Radiolarian

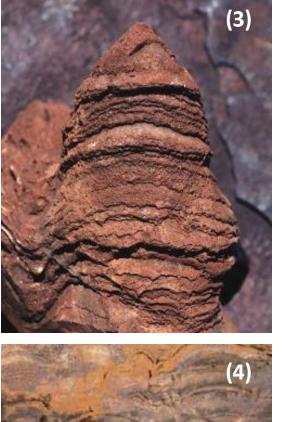
Photo: Sarah Spaulding

Diatoms





3.4-billion years stromatolite Strelley Pool Chert (Australia)



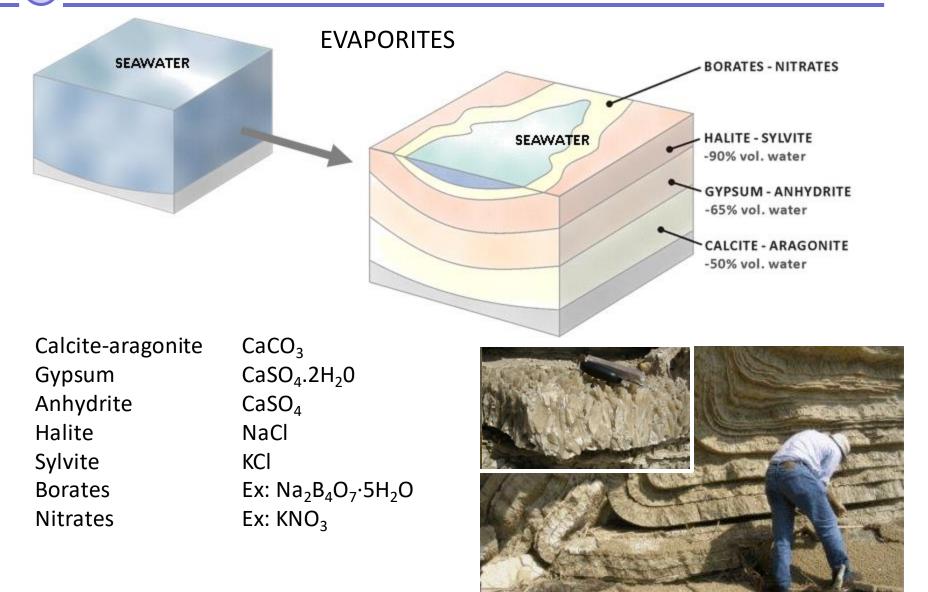
 $Ca^{2+} + 2HCO^{3-} \leftrightarrow CaCO_3 + CO_2 + H_2O$

 $6H_2O + 6CO_2 + sunlight \rightarrow C_6H_{12}O_6 + 6O_2$

Calcification

(1) Modern stromatolites at Shark Bay, Australia (P. Harrison, Wiki.) (3) Fossil stromatolite (K. McNamara, www.geolsoc.org.uk)
 (2) Modern stromatolite (http://phys.org)
 (4) Fossil stromatolite (Alwood et al. 2006, Nature)

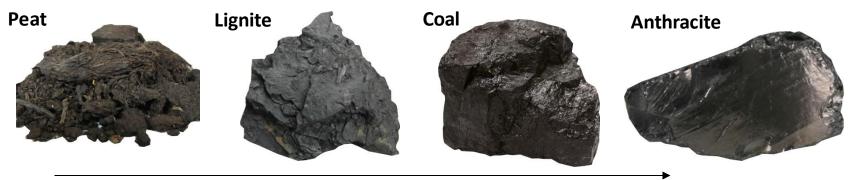
3. Chemical sedimentary rocks



Messinian (~5.5 Ma) evaporite composed of gypsum Source: F. Boulvain (University of Liege)

4. Organic sedimentary rocks

Note that biochemical and organic sedimentary rocks can be called biological sedimentary rocks



Burial (increasing P and T)

Another example: oil shale

- Problems of preservation:
 - 1. **RECYCLING** (organic matter in water column consumed by organisms)
 - 2. **OXIDATION** (bacterial and abiotic decay of organic matter)

 $\mathrm{C_6H_{12}O_6} + \mathrm{6O_2} \rightarrow \mathrm{6CO_2} + \mathrm{6H_2O}$

- Conditions of preservation:
 - 1. HIGH ACCUMULATION RATE (e.g. plant debris, micro-organisms)
 - 2. ANOXIA or LOW [O₂] (e.g. restricted water circulation)

Coal mine (Carboniferous, Graissessac, France)



F. Boulvain (ULg)

Different origin of coal and oil

 CH_4

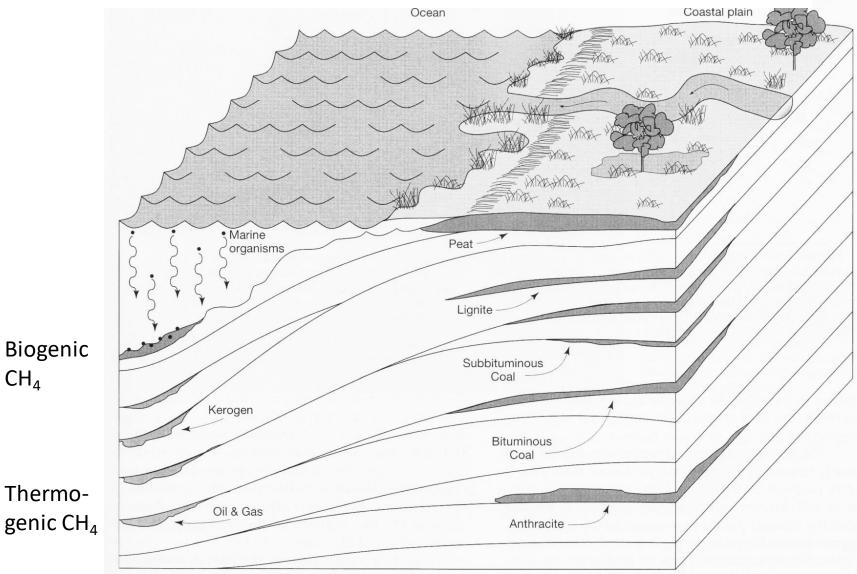


FIGURE 2.13 Schematic diagram illustrating how the burial of terrestrial organic matter can result in the formation of coal, and the burial of marine organic matter can result in the formation of oil and gas. The increasing temperature and pressure at greater depths of burial both compact and modify the terrestrial matter as it progresses through the various ranks of coal. Marine organic debris is converted into a waxy material called kerogen; upon additional heating, the kerogen is converted into petroleum.

TABLE 5.2 Representative compositions of living matter and fossil fuels

	Major Constituents (wt%)					
Substances	Lipids		Proteins		Carbohydrates	
Green plants	2		7	75		
Humus	6		10	77		
Phytoplankton	11		15		66	
Zooplankton	15		53		5	
Bacteria (veg.)	20		60	20		
Spores	50		8	42		
Part B: Petroleum		and the second	a de la composición d		The second	
	Elemental Composition (wt%)					
Substances	С	н	S	N	0	
Lipids	80	10			10	
Proteins	53	7	2	16	22	
Carbohydrates	44	6	<u> </u>		50	
Lignin	63	5	0.1	0.3	31	
Kerogen	79	6	5	2	8	
Natural gas	75–80	20-25	trace-0.2	trace-minor	_	
Asphalt	81–87	9-11	0.3–6	0.8-2.2	0–4	
Petroleum	82–87	12–15	0.15	0.1–5	0.1–2	
Part C: Coal	aft soft to	an Barran	(anigni) for			
Internet for the later of	Elemental Composition (wt%)					
Substances	С	Н	S	N	0	
Peat	21.0	8.3	3 100-00-00	1.1	62.9*	
Lignite	42.4	6.6	5 1.7	0.6	42.1	
Sub-bituminous	76.3	4.7	0.5	1.5	17.0	
Bituminous	87.0	5.4	1.0	1.4	5.2	
Semianthracite	92.2	3.8	3 0.6	1.2	2.2	
Anthracite	94.4	1.8	3 1.0	0.7	2.1	

Part A: Living Matter

*Remainder is ash and moisture.

(From Chilingarian and Yen. *Bitumens, Asphalts, and Tar Sands,* and from *Coal Development, S.* Bureau of Land Management, 1983.)